

## Regional and local ionospheric models based on Millstone Hill incoherent scatter radar data

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[1] Local and regional statistical models to describe Millstone Hill incoherent scatter radar observations of electron density, electron temperature and ion temperature since 1976 are developed using a bin-fit technique. The local models generate ionospheric variations with local time, day number, and altitude from 150–1000 km. The prior day's F107 and the Ap index from the previous 3 hour period are keyed inputs to specify solar and geomagnetic activity. The regional models have a latitude coverage of 32–55° geodetic and an altitude coverage of 200–600 km. These climatology models are capable of reproducing primary ionospheric variation features seen in previous studies as well as several newly revealed features, such as the semiannual variation of electron density. They are accessible through the World Wide Web at the URL <http://www.openmadrigal.org>. *INDEX TERMS*: 2447 Ionosphere: Modeling and forecasting; 6952 Radio Science: Radar atmospheric physics; 2443 Ionosphere: Midlatitude ionosphere; 6929 Radio Science: Ionospheric physics (2409)

### 1. Introduction

[2] Given that global models may smear out features which are unique to a particular region, regional and local models can be very useful. Millstone Hill (42.6°N, 288.5°E, Apex magnetic latitude 54°) is a unique midlatitude region in Eastern North America, and a region of critical importance for the United States National Space Weather Program (NSWP). At an L value of 3 it lies near the plasmapause boundary and may be considered “subauroral” during geomagnetically disturbed conditions. In this region the fact that the geomagnetic latitude is about 12° higher than the corresponding geographic latitude may cause high conductivities, and thermospheric circulations over this near-magnetic pole site may lead to interesting annual/semiannual ionospheric variations [Rishbeth, 1998; Rishbeth et al., 2000]. Since the 1960s, incoherent scatter (IS) measurements of electron density  $N_e$ , electron temperature  $T_e$ , ion temperature  $T_i$ , and line-of-sight ion drifts have been acquired over Millstone Hill. The favorable location at subauroral latitudes combined with the great operational range afforded by the radar's steerable antenna permit observations over a latitude span encompassing the region between the polar cap and the near-equatorial ionosphere.

[3] From the extensive database of Millstone Hill IS observations, empirical models of basic and derived IS parameters, including  $N_e$ ,  $T_e$ ,  $T_i$ , electric fields and parallel drifts, are being developed. Here we present results for the three scalar parameters. These models are very important in validating theoretical models and empirical global models such as International Reference Ionosphere (IRI) [Bilitza, 2001], and also they can be applied to addressing outstanding scientific issues related to the ionosphere and thermosphere climatology: e.g., the annual and solar cycle

variations of  $N_e$ , the ionospheric trough, the so-called dusk effect, the dependence of  $T_e$  and  $T_i$  on magnetic activity, etc. In this paper, the data used cover the 25-year period from February 1976 to August 2001, over a 200–600 km height range and a 32–55° geodetic latitude span for regional modeling, and over a 150–1000 km height range and 39–45° latitude span for the higher-resolution local modeling. Most of the measurements used in the regional model were within a few degrees of Millstone Hill's longitude, though at high altitudes some were as much as 25° to the west. At 1000 km, the data used in the local model can be up to 7° to the east or west of Millstone. We employ a bin fit technology to sort the data and construct models. This will be described, and then we discuss the distribution of the sorted data bins. Next, we give general descriptions of the main features of our model results. Finally, we summarize this work and provide information on model availability.

### 2. Data and Method

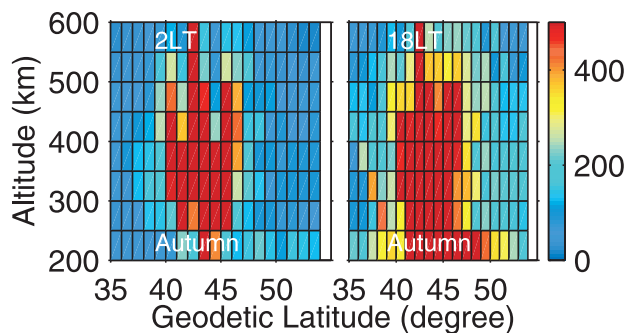
[4] The Millstone Hill UHF IS radar system operates with a zenith-directed 68 m diameter fixed parabolic antenna, which commenced operation in 1963, and a fully-steerable 46 m antenna, which commenced operation in 1978. The electron density and plasma temperatures are determined from the received power and spectrum. Experiments are carried out for more than 1000 hours per year. Data are archived in the Millstone Hill Madrigal online database system, which contains an extensive body of ground-based measurements and models of the Earth's upper atmosphere and ionosphere. Altogether, 907 experiments were included in the models.

[5] The statistical models are computed using a bin-fit technique, which combines data binning with least squares fitting. The measured  $N_e$ ,  $T_e$  and  $T_i$  are separated into bins according to local time, altitude, and the day of year (to represent seasonal variations), and also to geodetic latitude for regional modeling. The bin size is 1 hour for local time and 61 days for the seasonal variation. For the regional model, the bin size is 1° from 32–55° geographic latitude, and 50 km from 200–600 km. For the local model, there are 12 altitude bins centered at 150, 175, 200, 225, 250, 300, 350, 400, 500, 600, 800, and 1000 km, and data are limited to 39–45° geodetic latitude. In each bin, the dependencies on solar and magnetic activity are determined through a sequential least squares fit based on Givens transforms [Gentleman, 1973] to the following equation:

$$P = \beta_0 + \beta_1 \times f + \beta_2 \times a + \beta_3 \times f \times a$$

where  $P$  is either  $N_e$ ,  $T_e$  or  $T_i$ , the  $\beta$ s are fitting coefficients, and  $f = (F107 - 135)/135$  and  $a = (Ap - 15)/15$  are the normalized F107 and Ap indices. F107 is the previous day's 10.7 cm solar flux index, and Ap is the 3-hourly equivalent range index  $A_p$  for the previous 3 hours. Deviations of actual data from the model represent the remaining day-to-day variability due to such causes as tidal forcings, gravity waves, uncertainties in the solar EUV flux and high latitude forcings which are difficult to accurately specify. While the physics involved in ionospheric responses to F107 and Ap is rather complicated, we opt for the linear approximation for simplification while giving the correct general trends of variations

<sup>1</sup>Deceased 20 October 1999.



**Figure 1.** Distribution of number of data points with geodetic latitude and altitude for the regional models. The cell color represents the data number for the lower-left corner grid.

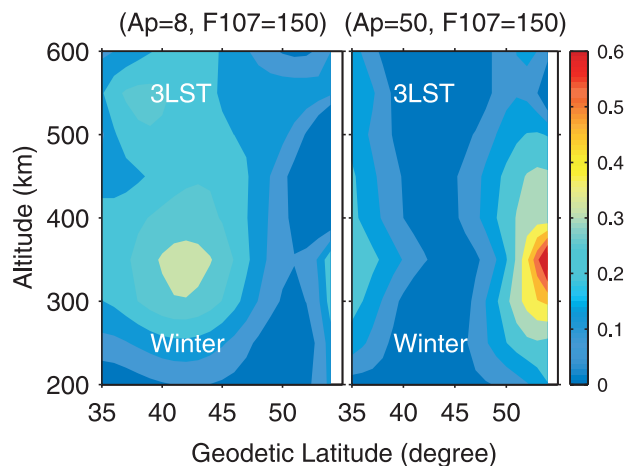
for our climatology models. A separate study demonstrating detailed variation trends of the data with F107 and Ap indicates the general validity of our linear approximation in the fit formula.

### 3. Data Distribution

[6] The data distributions of the electron density and plasma temperatures are essentially the same. Therefore we discuss only the  $N_e$  data. For local modeling, the data density is very high in the 250–500 km range, at about 70–100 counts/km for a given day number and local time bin, corresponding to, for example, about 3000–7000 data points for the 300 km bin. For lower or higher altitudes, the number of data points is on the order of 1000 for daytime and 500 for nighttime. More data were obtained in the January to April and September to October periods than in other months. For the regional models, more than 350 data points are present during the day, and 200 during the night in the latitude range 38–50° as shown in Figure 1. Most data are from periods of low magnetic activity, with  $A_p < 20$ , and there are also many data for  $20 < A_p < 40$ .

### 4. Result

[7] Our local models reproduce the occurrence and evolution of a few interesting midlatitude ionospheric phenomena, including the well-known “seasonal anomaly” where the midday  $N_e$  is higher in winter than in summer and the morning  $N_e$  increases more steeply in winter [Rishbeth and Setty, 1961] due to the  $O/N_2$  effect, the midday density depression (“bite-out”) which relates to the northward neutral wind drag [Kohl and King, 1967], the evening density peak in summer [Evans, 1965] (see the left panel of Figure 2), the “predawn effect” in the electron temperature  $T_e$  [Carlson and Weill, 1966],  $T_e$  enhancements in the morning and in the afternoon [Brace and Theis, 1981], and the unusual high  $T_e$  on winter nights [Evans, 1973]. There appears a striking feature that has not

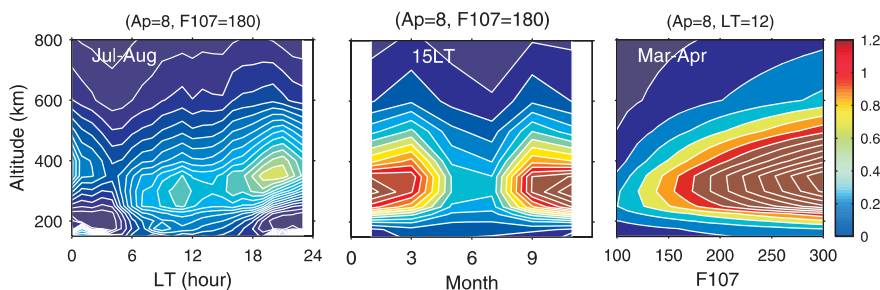


**Figure 3.** Subauroral ionospheric trough at 0300 LT in winter given by the regional electron density model for F107 = 150 and  $A_p = 8$  (left) and  $A_p = 50$  (right). Electron density is in units of  $10^{12} \text{ m}^{-3}$ .

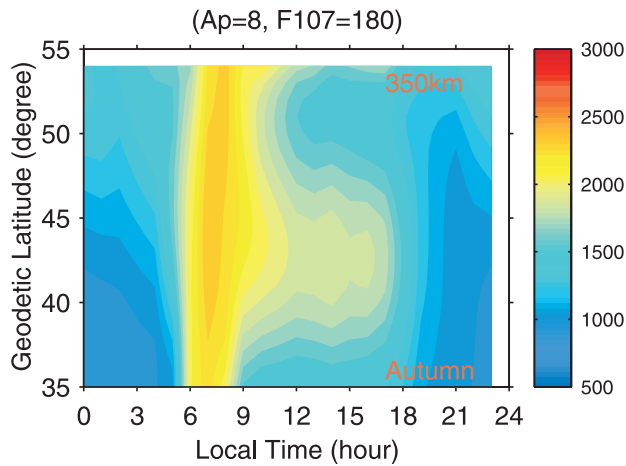
previously been reported for Millstone Hill in a statistical manner: the semiannual variation of  $N_e$ , most pronounced above the  $F_2$  peak, as shown in the center panel of Figure 2. During the day,  $N_e$  increases from winter, reaching an maximum in spring, then decreases toward summer when it reaches a minimum; a similar pattern follows in the second half of the year. Relative to other midlatitude sites where semiannual variations occur [Balan et al., 1998], Millstone Hill is nearer the north geomagnetic pole, and the winter density is supposed, by the well-established thermospheric circulation theory [Rishbeth et al., 2000], to be highest, exceeding the equinox density when the zenith angle effect is less important. This semiannual feature needs to be further investigated.

[8] As for the solar activity dependence,  $N_e$  and  $T_i$  generally increase with F107, but the  $N_e$  increase becomes smaller for very large F107 in winter during noon and later [Balan et al., 1994]. Figure 2 (the right panel) shows an example of electron density responses to the F107 change.  $T_e$  responses are more complicated depending on local time, season and altitude [see also Bilitza and Hoegy, 1990]. It generally increases with F107, but decreases in summer when  $N_e$  is not very high. With increasing geomagnetic activity,  $N_e$  is typically reduced in summer and equinox periods (negative storm),  $T_e$  is elevated during the night, and  $T_i$  is elevated during almost the entire day.

[9] The midlatitude ionospheric trough is often seen in Millstone Hill radar data [Holt et al., 1983]. This feature is clearly seen in the regional  $N_e$  model (Figure 3) for high magnetic activity. The trough is basically an afternoon and nighttime feature, though it may also appear during the day [Evans et al., 1983; Vo and Foster, 2001]. This ionospheric main trough is considered to be caused by



**Figure 2.** Iso-density contours with local time and altitude in summer (left panel), with month and altitude (middle panel), and with F107 and altitude (right panel) computed with the local electron density model for  $A_p = 8$ . Electron density is in units of  $10^{12} \text{ m}^{-3}$ .



**Figure 4.** Electron temperature contours with geodetic latitude and local time for altitude 350 km in autumn with  $F107 = 180$  and  $A_p = 8$ .

plasma convection through the nightside in the absence of ionisation sources and  $O^+$  chemical recombination in the presence of large electric fields [Rodger et al., 1992; Foster, 1993]. Detailed discussions of the trough structure seen near Millstone Hill have recently been given by Vo and Foster [2001]. Associated with the density depression in the trough,  $T_e$  is found to be high due to the reduced cooling rate. Compared with IRI-95, the seasonally averaged  $N_e$  above Millstone Hill is lower. Also significantly different from the IRI is the occurrence of a  $T_e$  morning enhancement in winter: it peaks sharply near 0800 LT, while it does not seem to appear during winter in the IRI.  $T_e$  typically increases with increasing magnetic activity, especially at higher latitudes and at night. An obvious latitudinal variation of  $T_e$  is the decrease toward low latitudes during the night. It is most pronounced during summer, and is also present in spring and autumn. The  $T_e$  decrease does appear during winter nights, but is not as evident as in other seasons. Figure 4 shows  $T_e$  in autumn for 350 km. This feature may be ascribed to the latitudinal difference of the heat flow from the topside ionosphere [Schunk and Nagy, 2000].

## 5. Conclusion

[10] Local and regional statistical models describing Millstone Hill IS radar observations of electron density, electron temperature and ion temperature have been developed using a bin-fit technique. The local models generate ionospheric variations with local time, day number, and altitude over 150–1000 km, as a function of the previous day's F107 and the  $A_p$  index for the previous 3 hours. F107 and  $A_p$  are keyed inputs to specify solar and geomagnetic activity. The regional models have a latitude coverage of 32–55° geodetic for 200–600 km altitude range. These climatology models are capable of reproducing primary ionospheric variation features seen in previous studies as well as several newly revealed features, such as the semiannual variation of electron density. They are accessible through the World Wide Web at the URL <http://www.openmadrigal.org>. The user can select model type (regional or local model) and parameter type ( $N_e$ ,  $T_e$ , or  $T_i$ ), and specify F107

and  $A_p$  indices to get outputs in the form of various contour plots as well as ASCII data ready to be sent to the user's E-mail address. Several animated pictures demonstrating dynamic changes of ionospheric parameters are also available. We also provide portable software for recovery of model values.

[11] **Acknowledgments.** We thank the members of the Haystack Observatory Atmospheric Sciences Group for assembling and maintaining the Madrigal Database of Millstone Hill incoherent scatter radar observations, which were the basis of this study. This research was supported by NSF Space Weather Grant ATM-9819413. The Millstone Hill incoherent scatter radar is supported by a cooperative agreement between the National Science Foundation and the Massachusetts Institute of Technology.

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